

Proyecciones hacia finales de siglo del nivel del mar extremo sobre la base del escenario de emisiones RCP 8.5

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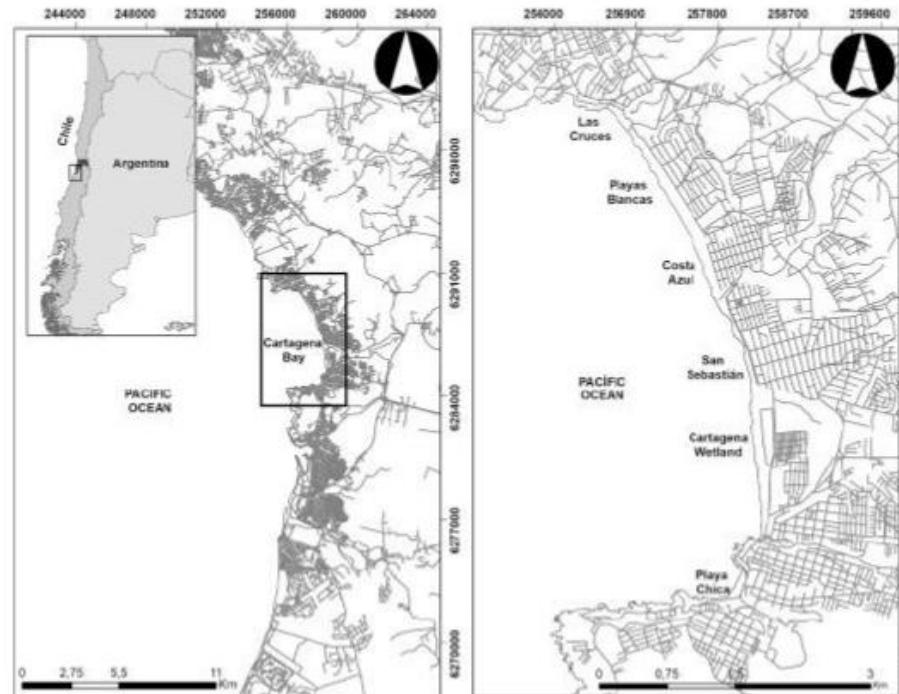
INTRODUCCIÓN

SECTOR DE ESTUDIO - CARTAGENA BAY (33.5°S)



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Cartagena is a Chilean commune located in the San Antonio Province, Valparaíso Region. The commune spans an area of 245.9 km². Is one of the largest in central Chile and currently one of the most anthropized by urbanization. Its physical-natural attributes and history of occupancy allow its risk conditions, including the following, to be contextualized:



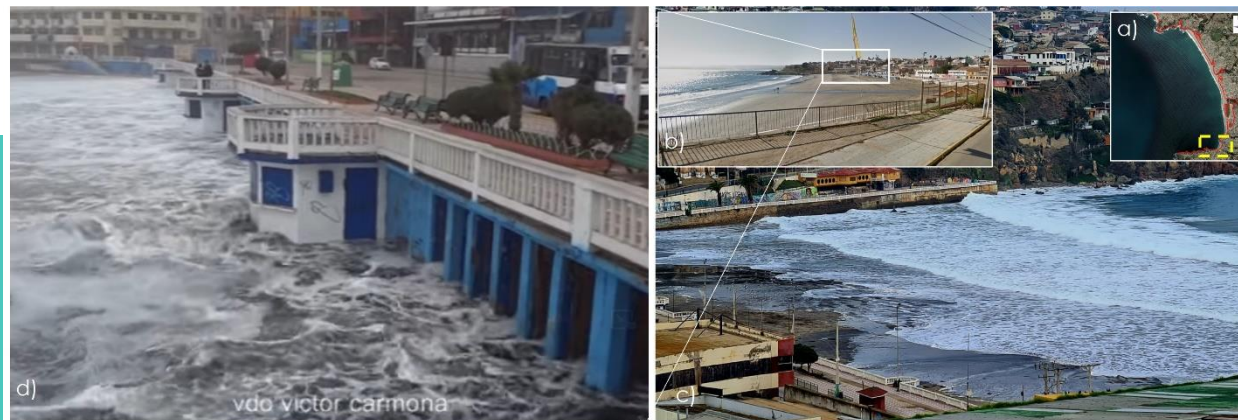
Source: Martinez et al. 2020

MOTIVACIÓN

INUNDACIONES COSTERAS EN CARTAGENA



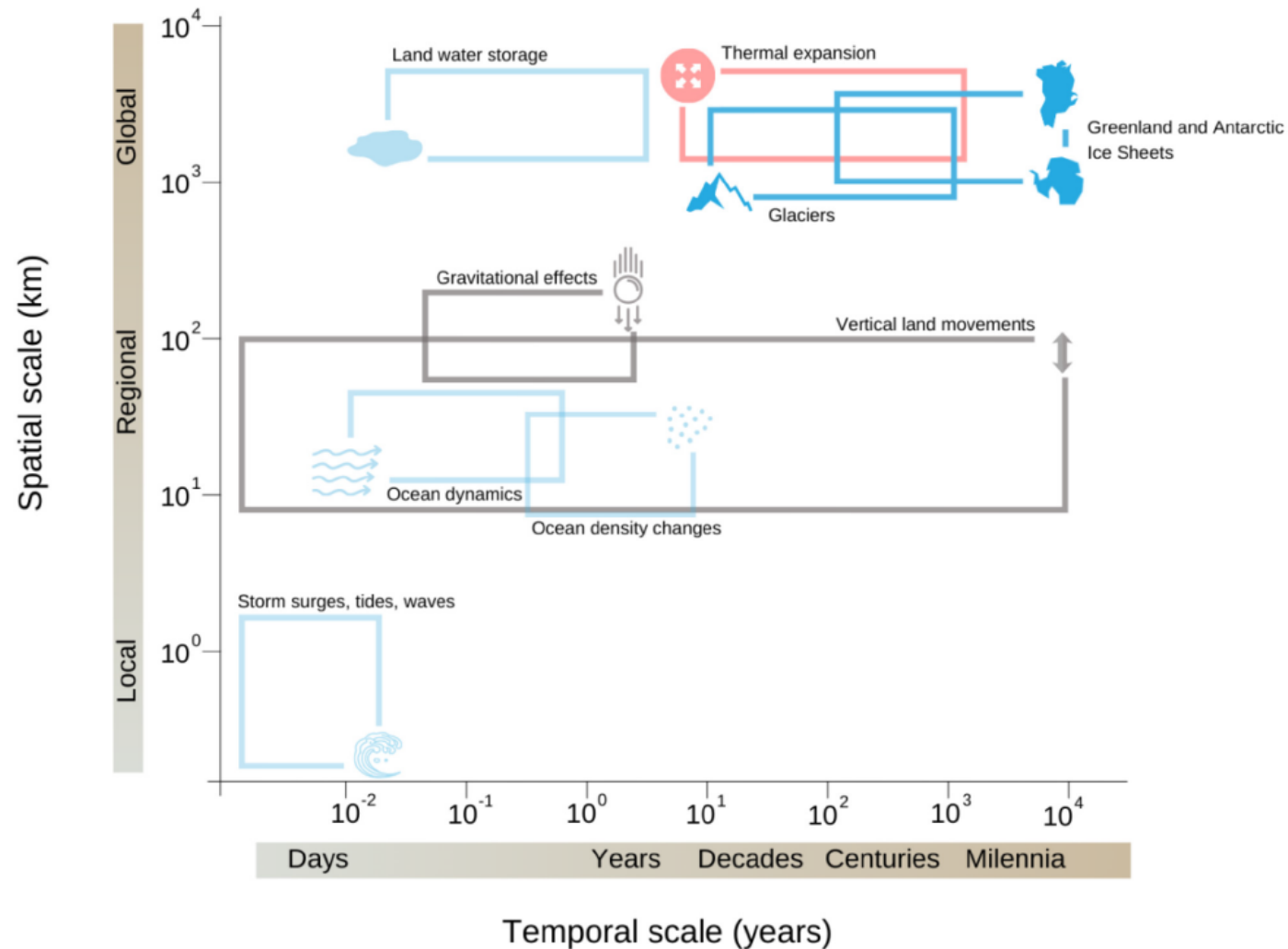
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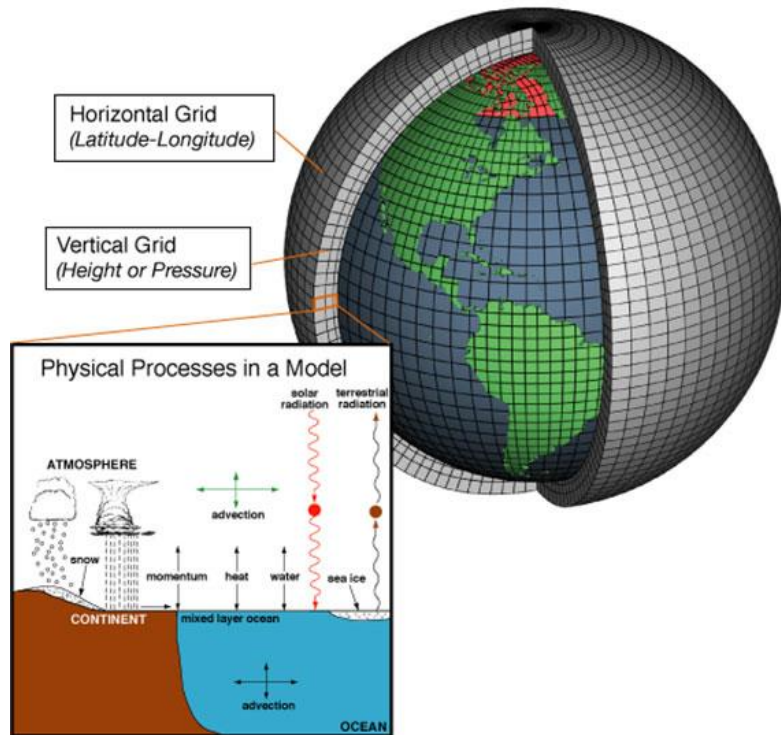
Marejada 21-22 Agosto 2021 Humedal de Cartagena. Fuente: P. Cortez (IMCartagena).

NIVEL DEL MAR EXTREMO

MÚLTIPLES ESCALAS



GLOBAL CLIMATE MODELS (GCM)



GCM [proyección]	Institución	Resolución Atmosférica Lon x Lat	Resolución temporal (viento) x (hielo marino)	Ensemble
ACCESS 1.0 historical: 1985-2004 midRCP85: 2026-2045 endRCP85: 2081-2100	CSIRO-BOM	1.88 × 1.25	3hrs x 1 día	r1i1p1
MRI-CGM3 historical: 1985-2004 midRCP85: 2026-2045 endRCP85: 2081-2100	Meteorological Research Institute	1.1 × 1.1	3hrs x 1 día	r1i1p1
MIROC5 historical: 1985-2004 midRCP85: 2026-2045 endRCP85: 2081-2100	Atmosphere and Ocean Research Institute, The University of Tokyo	1.4 × 1.4	3hrs x 1 día	r1i1p1
HadGEM2-ES historical: 1985-2004 midRCP85: 2026-2045 endRCP85: 2081-2100	Met Office Hadley Centre UK	1.88 × 1.25	3hrs x 1 día	r1i1p1
CMCC-CM historical: 1985-2004 midRCP85: 2026-2045 endRCP85: 2081-2100	Centro Euro- Mediterraneo per i Cambiamenti	0.7484 × 0.75	3hrs x 1 día	r1i1p1
EC-EARTH historical: 1985-2004 midRCP85: 2026-2045 endRCP85: 2081-2100	European Consortium	1.1215 × 1.125	3hrs x 1 mes	r8i1p1

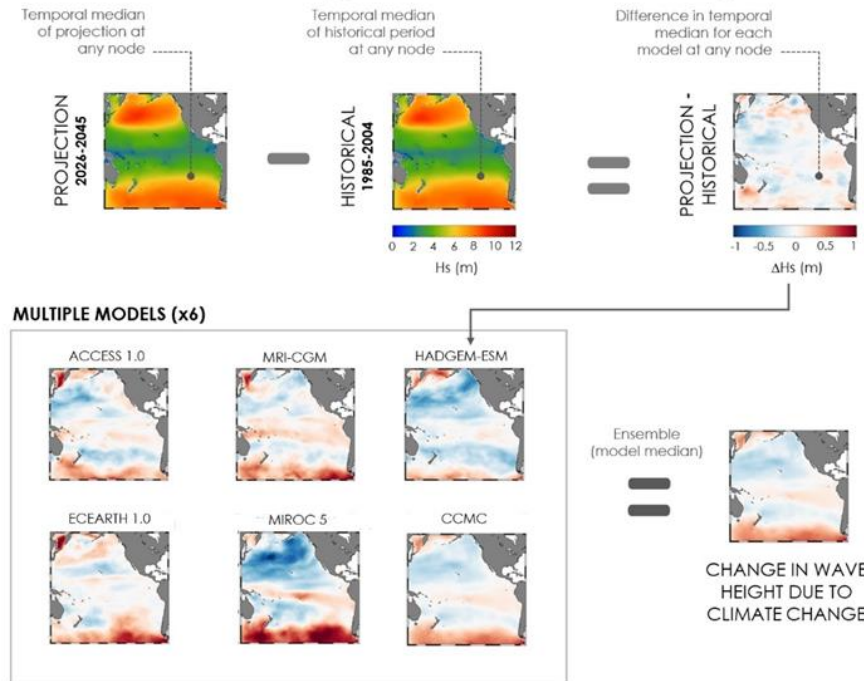
Hosted by  Department of Energy
Lawrence Livermore National Laboratory



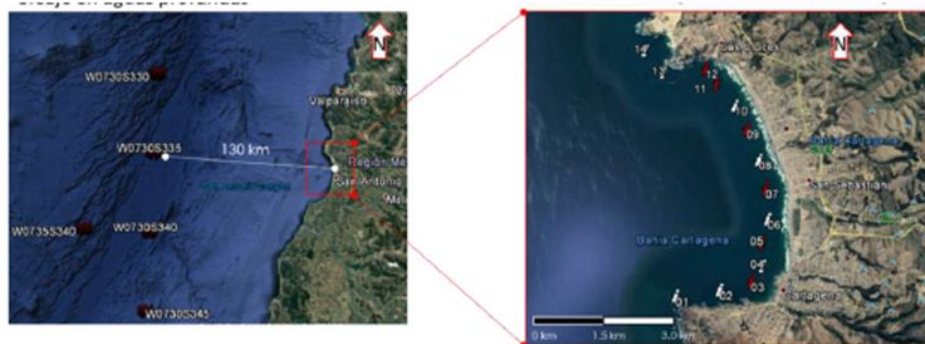
NLHPC
National Laboratory
for High Performance
Computing
Chile

CMIP5 involves 20 climate modeling research groups around the world with 40 GCMs. CMIP5 outputs

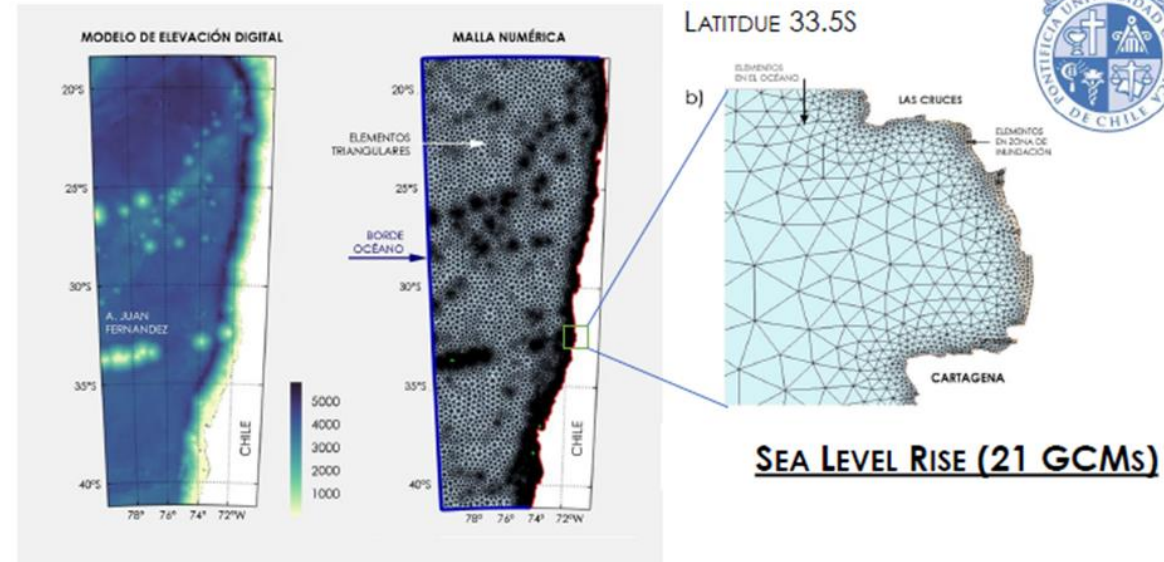
MULTIMODEL ENSEMBLE OF SIX GCM – SIGNIFICANT WAVE HEIGHT (WAVEWATCH III v6.07)



COASTAL WAVE PROPAGATION NODES (STWAVE 6.0)



ADCIRC v55

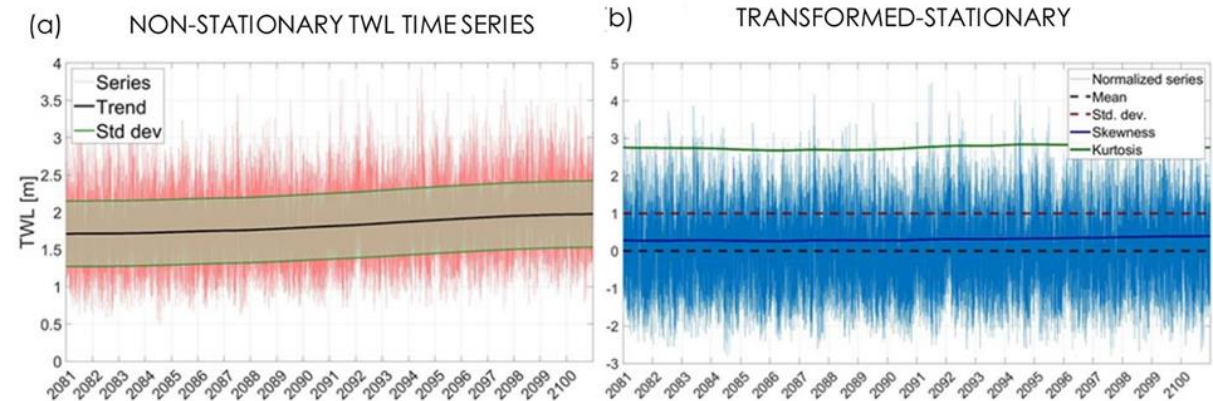


Additive linear model

$$TWL(t) = h_{runup}(t) + h_{storm\ surge}(t) + h_{sea\ level\ rise}(yearly) + h_{astro}$$

TOTAL WATER LEVEL TWL TIME SERIES

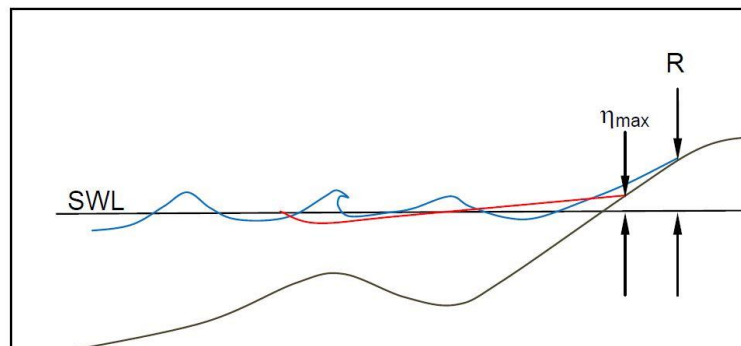
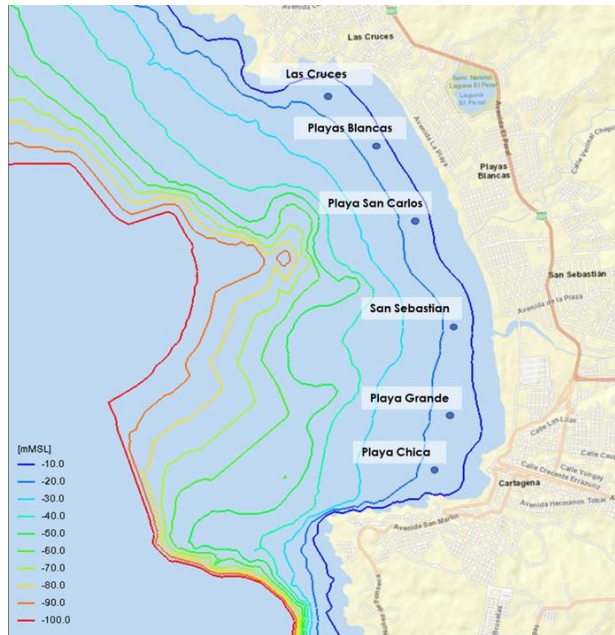
RUN-UP
STOCKDON (2006)



COASTAL RUNUP - STOCKDON ET AL. (2006)



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Beach	Transect	Transect slope	Slope used
Playa Chica	1	0.040	0.050
	2	0.060	
Playa Grande	3	0.060	0.070
	4	0.070	
	5	0.070	
	6	0.080	
Playa San Sebastian (wetland)	7	0.085	0.0850
	8	0.085	
	9	0.085	
	10	0.085	
Playa San Carlos	10	0.085	0.0913
	11	0.125	
	12	0.070	
	13	0.085	
Playas Blancas	13	0.085	0.0875
	14	0.110	
	15	0.085	
	18	0.070	
Playa Las Cruces	15	0.085	0.1033
	16	0.115	
	17	0.110	

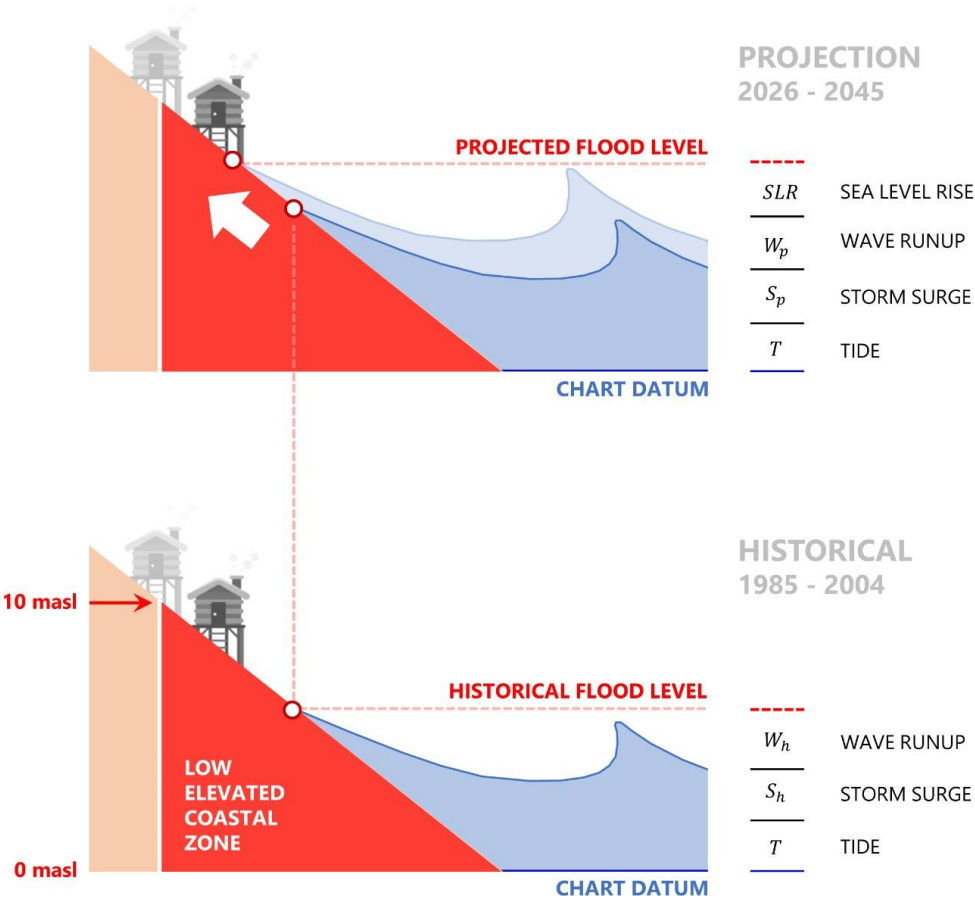
$$R_{2\%} = 1.1 \left[0.35 \tan \beta_f (H_0 L_0)^{1/2} + \frac{H_0 L_0 (0.5625 \tan^2 \beta_f + 0.0004)^{1/2}}{2} \right]$$

METODOLOGÍA

TOTAL WATER LEVEL



$$TWL(t) = h_{runup}(t) + h_{storm\ surge}(t) + h_{sea\ level\ rise}$$

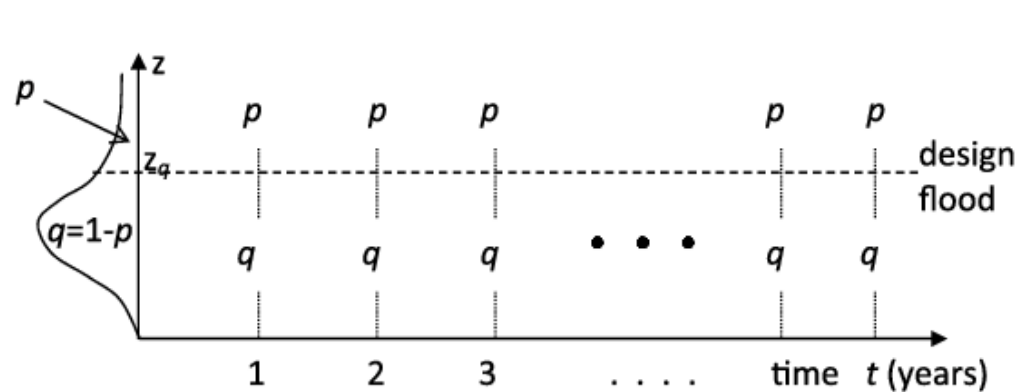


METODOLOGÍA

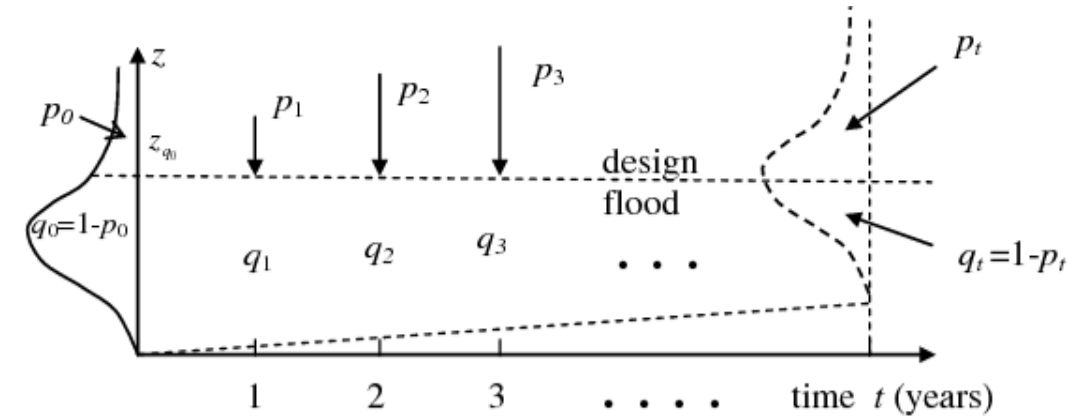
EXTREME VALUE – TRANSFORMED STATIONARY METHOD (MENTASCHI ET AL. 2016)



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Source: Salas and Obeysekera (2014)

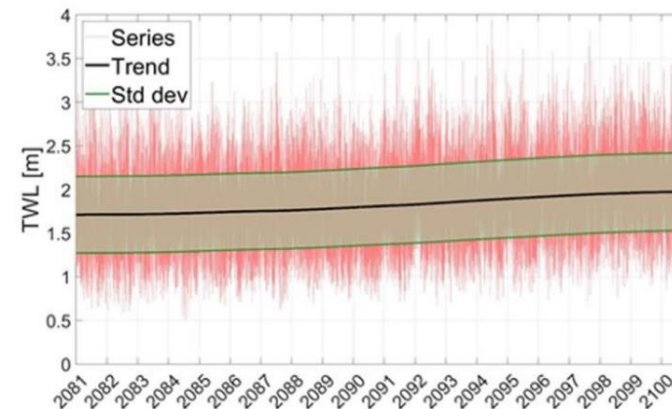


Normalize in time (remove non stationarity)

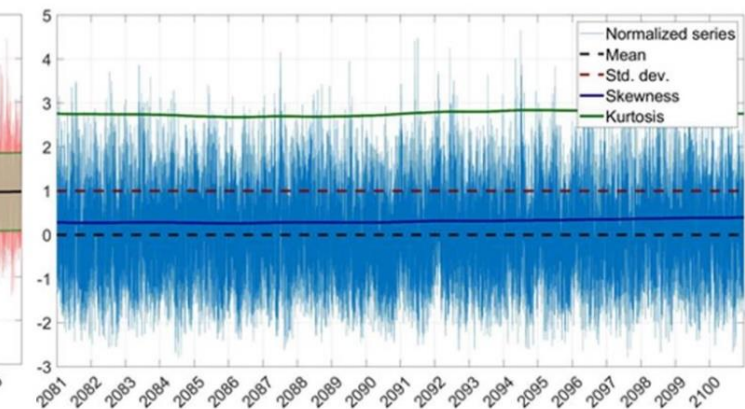
$$x(t) = \frac{y(t) - tr_y(t)}{std_y(t)}$$

For the ESL we use the method of peak over threshold (POT) and the generalized Pareto distribution (Pickands, 1975).

(a) NON-STATIONARY TWL TIME SERIES

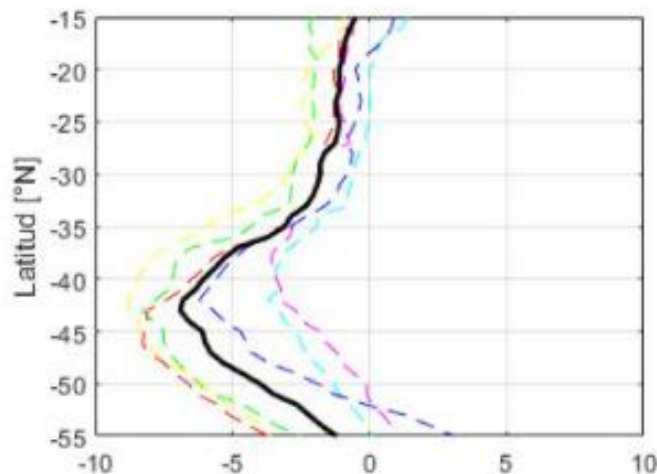


(b) TRANSFORMED-STATIONARY

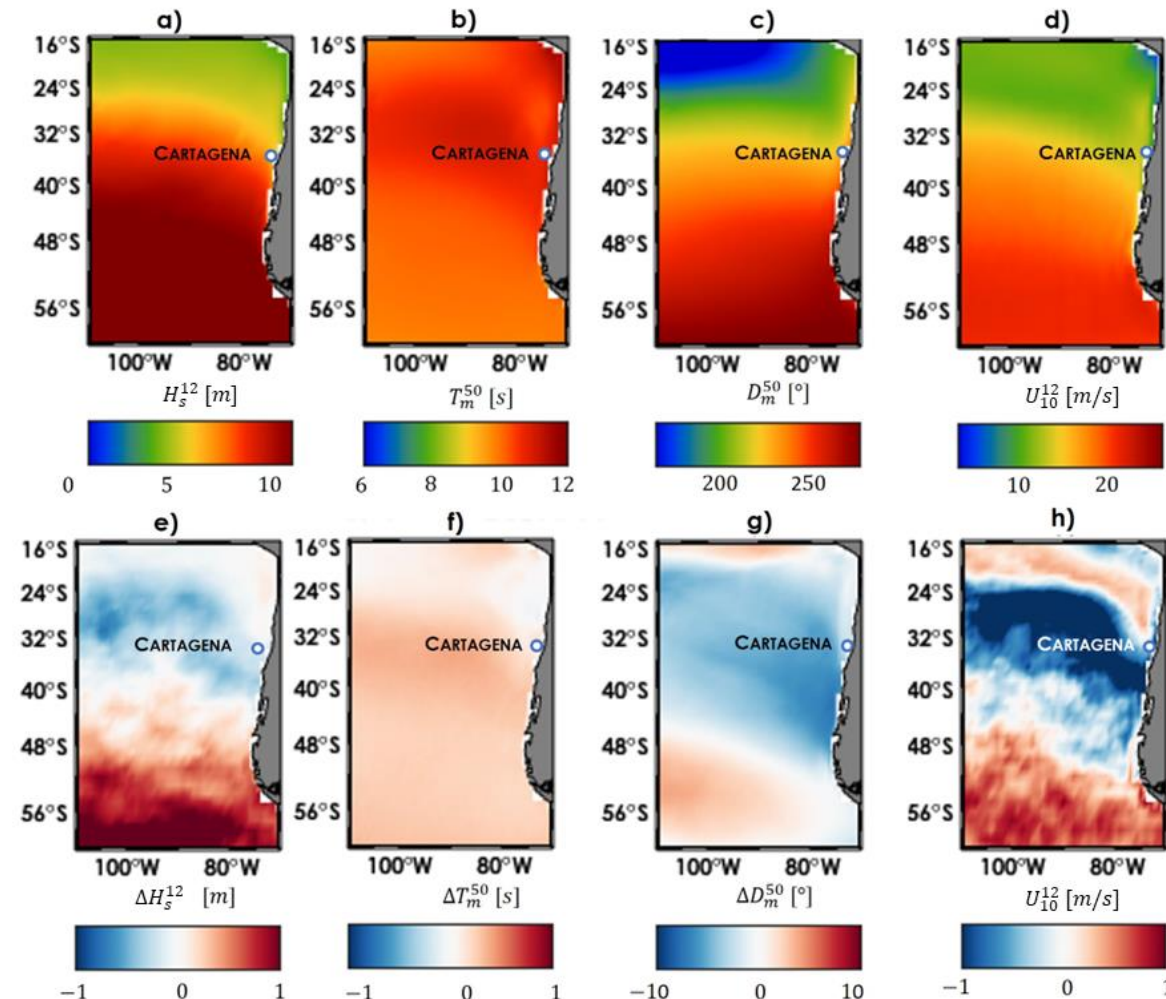


RESULTS - END CENTURY WAVE PROJECTIONS (2081-2100)

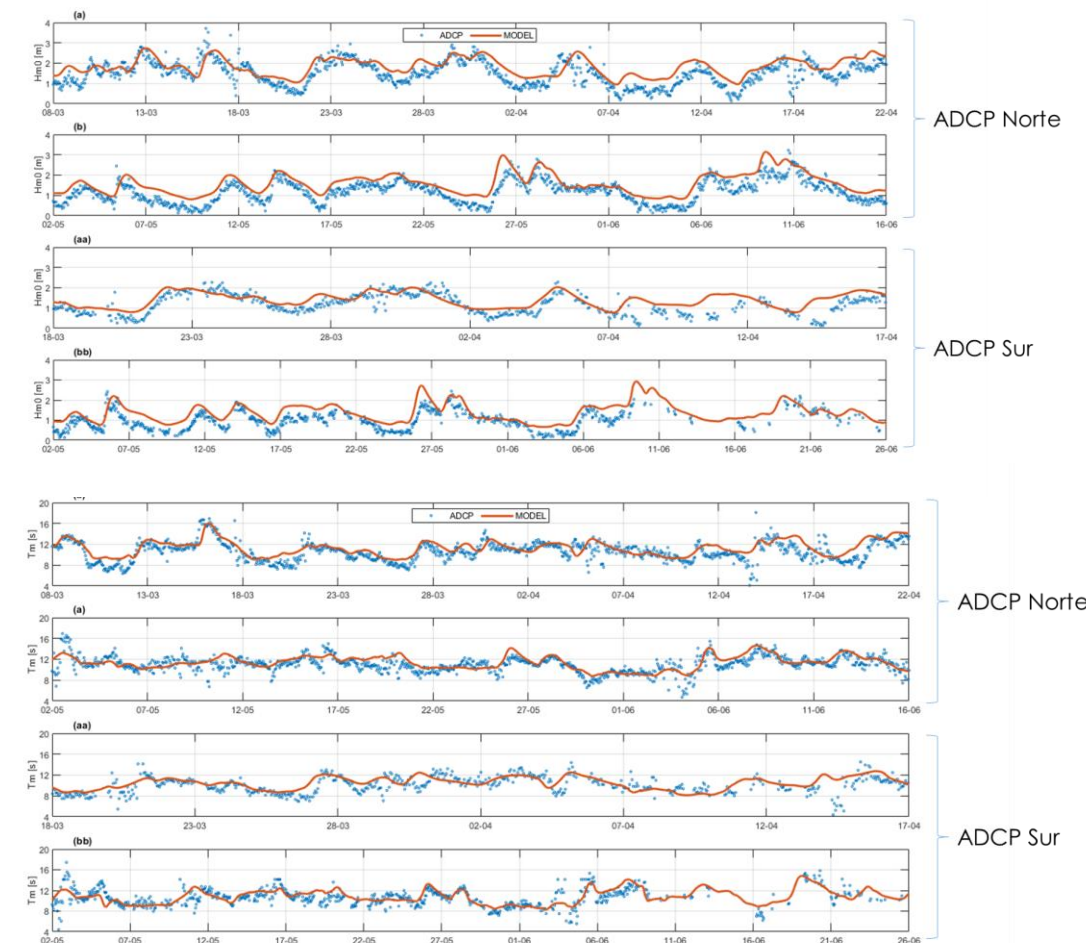
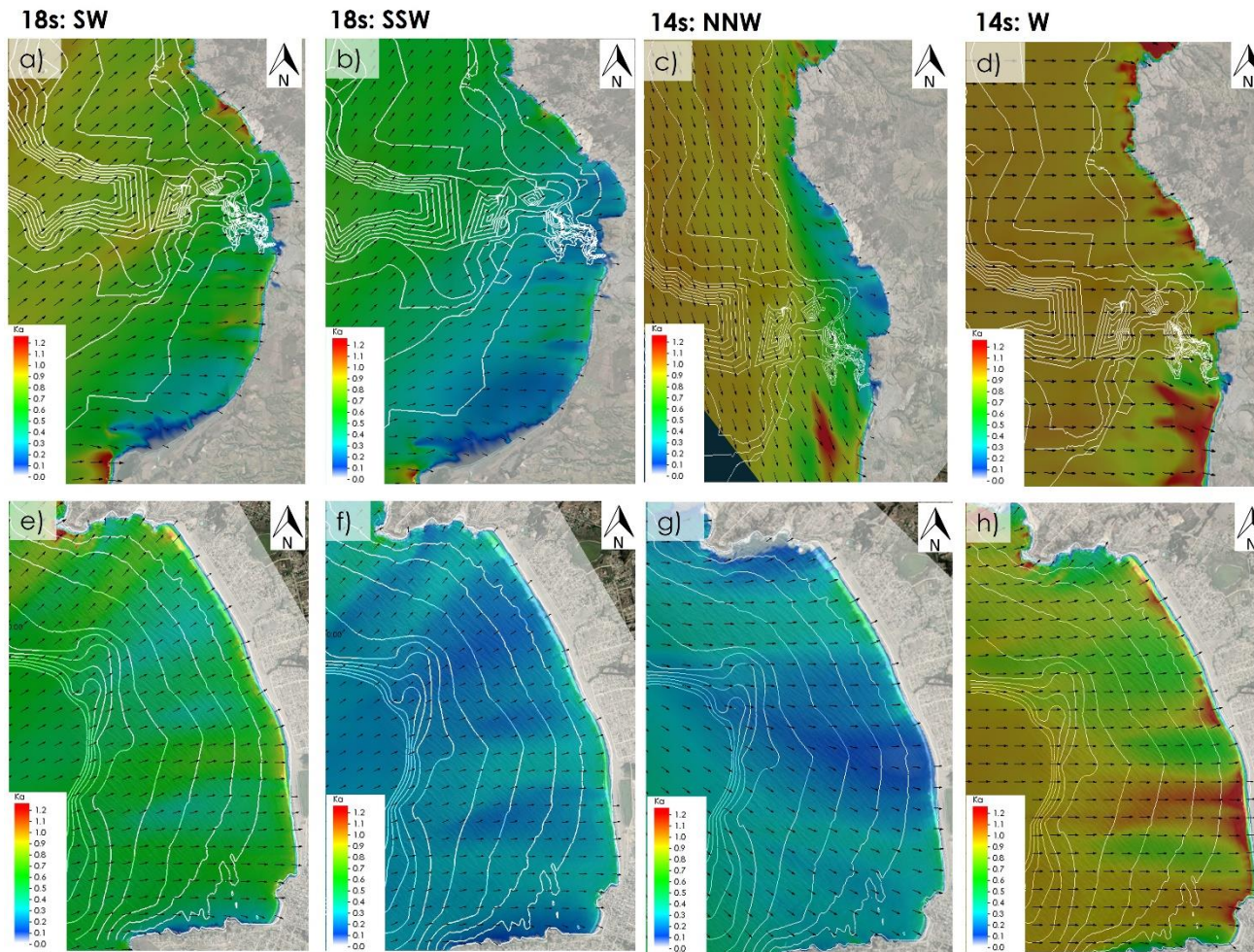
- Significant wave height yearly 12hrs exceedance: decreasing in central Chile (30-35°S). Increasing above latitudes 50°S
- Mean wave period 50th percentile: increasing in all Chilean coasts.
- Mean wave directions: counter clockwise shift



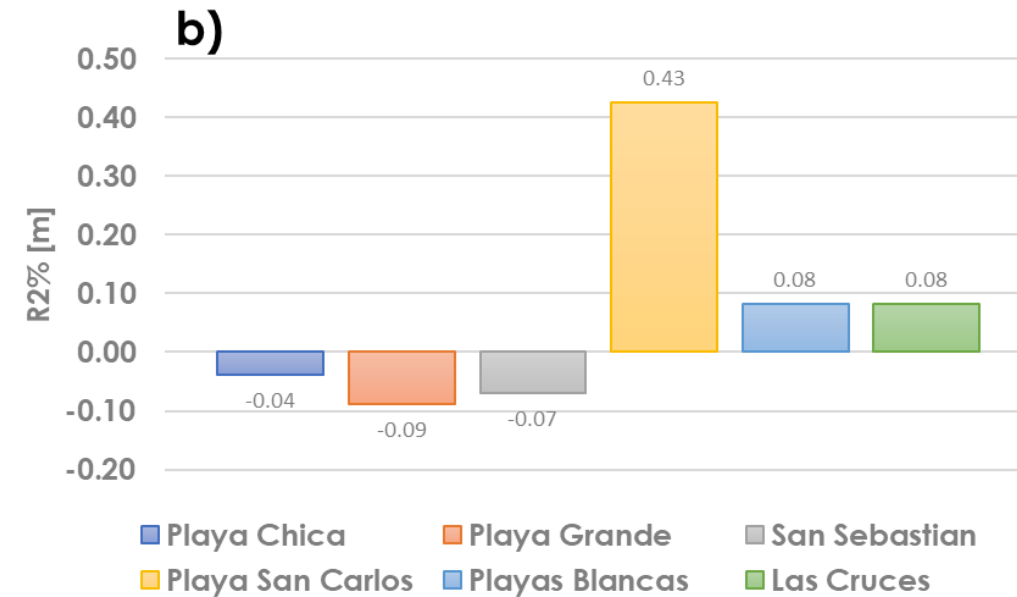
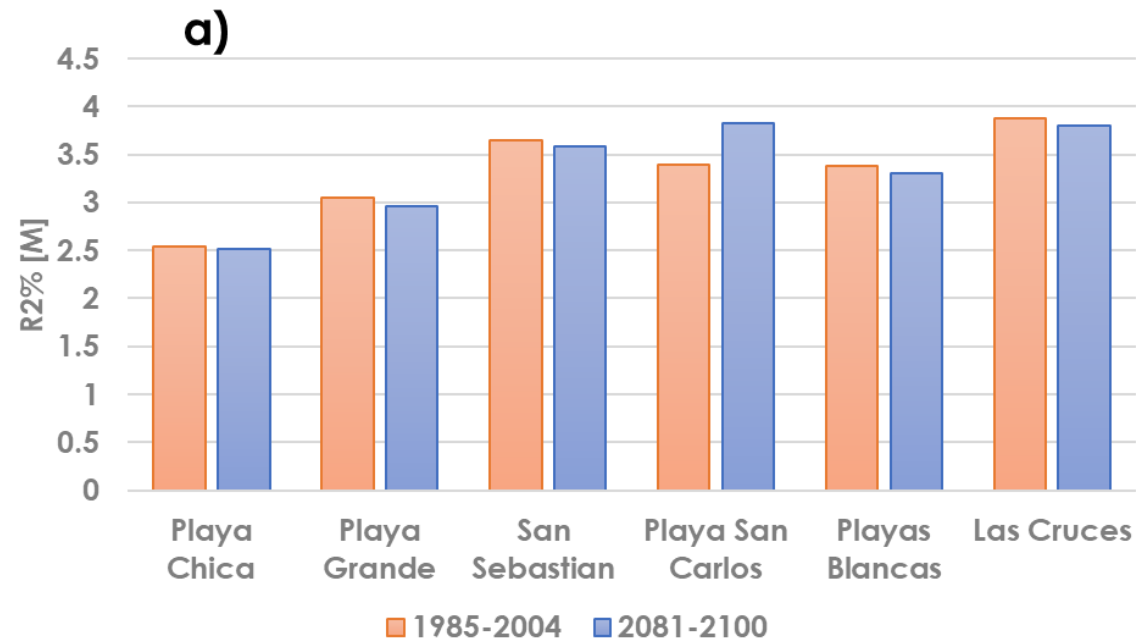
Projected - historical 50th percentile mean wave direction [°]



WAVE DOWNSCALING VALIDATION



RESULTS – 100 YEAR RETURN LEVEL WAVE RUNUP PROJECTIONS (2081-2100)



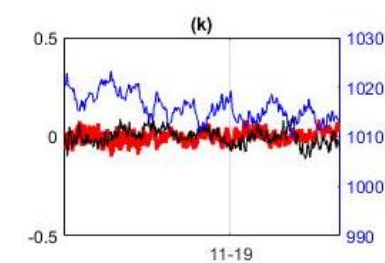
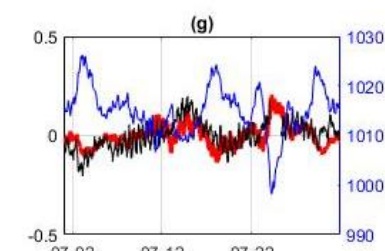
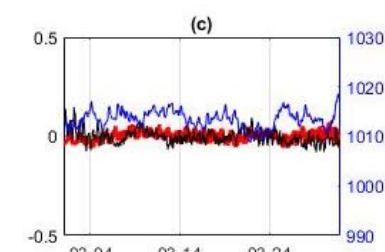
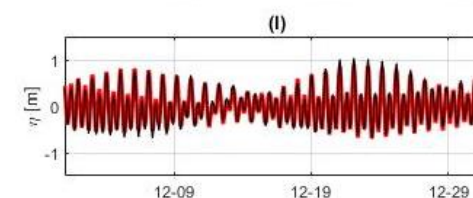
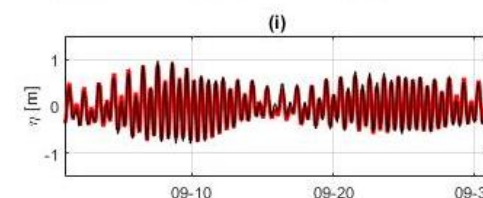
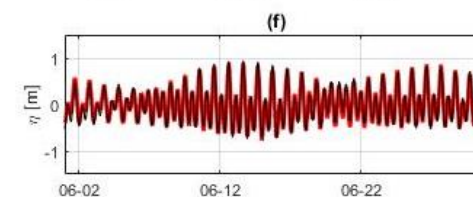
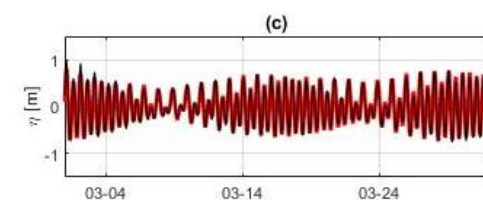
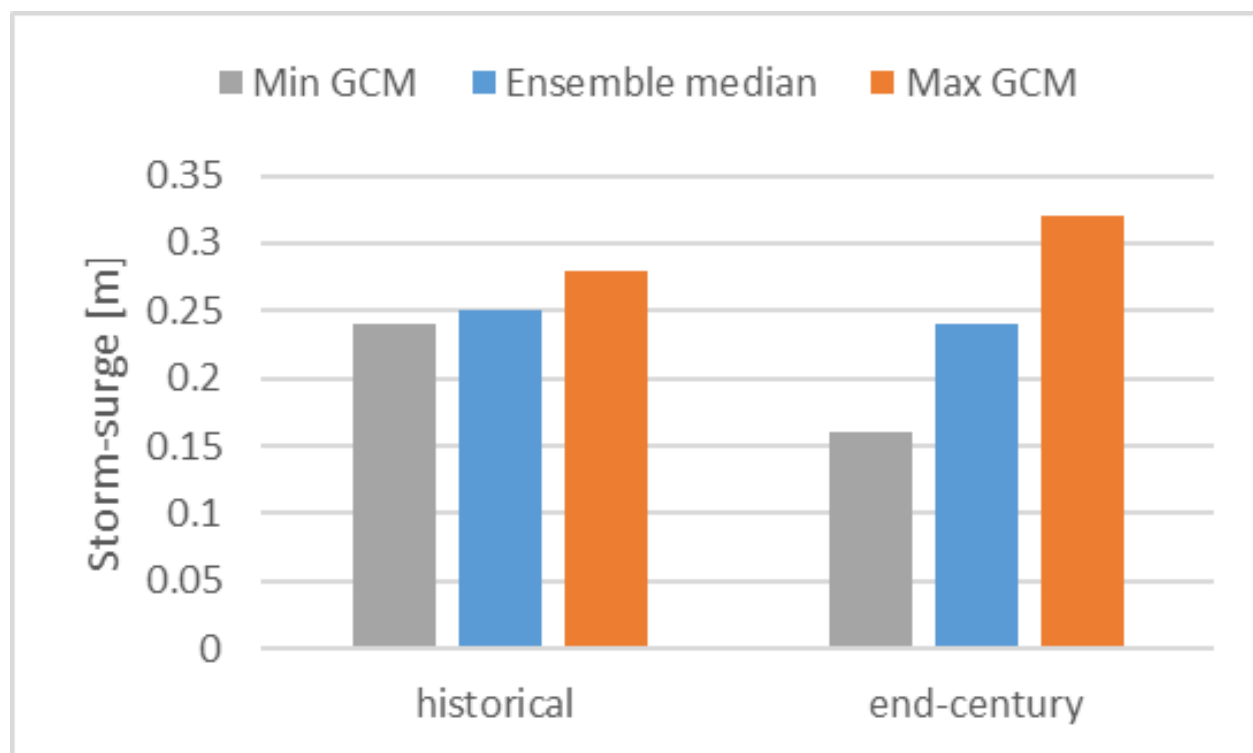
$$R_{2\%} = 1.1 \left[0.35 \tan \beta_f (H_0 L_0)^{1/2} + \frac{H_0 L_0 (0.5625 \tan^2 \beta_f + 0.0004)^{1/2}}{2} \right]$$

STORM-SURGE 100 YEAR RETURN LEVEL MODELO ADCIRC v55

ADCIRC VS SAN ANTONIO TIDAL GAUGE

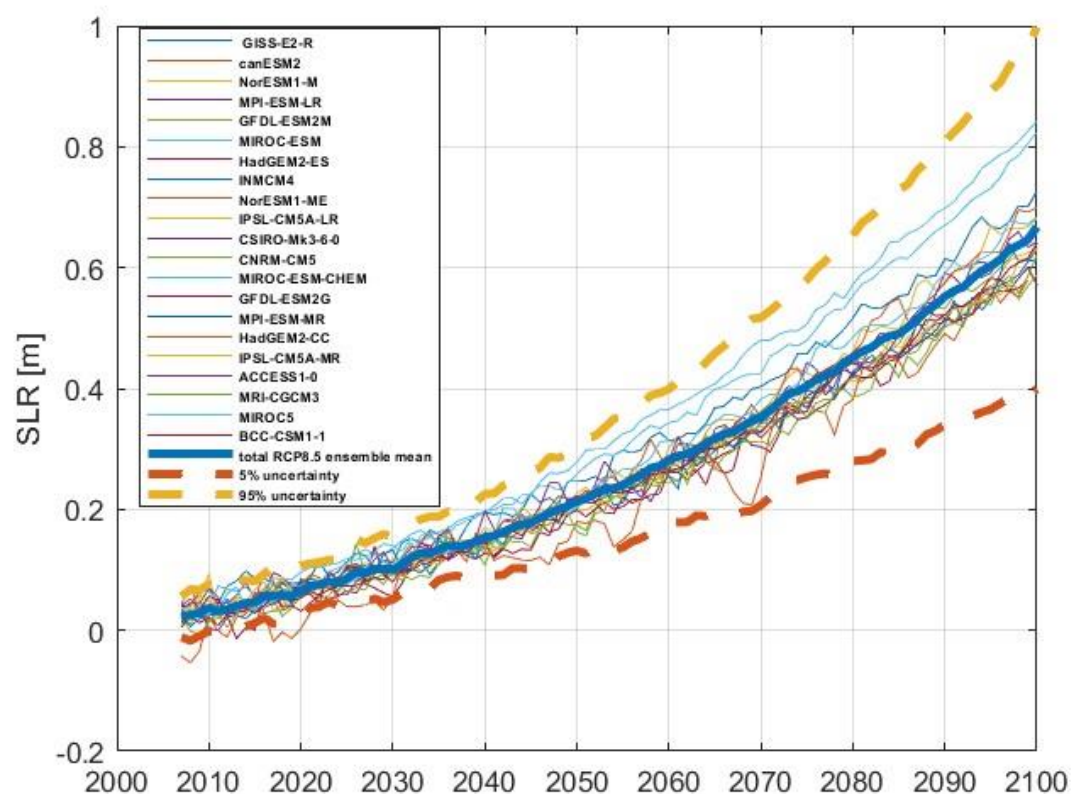
ASTRONOMICAL TIDES

STORM-SURGE: RESIDUAL



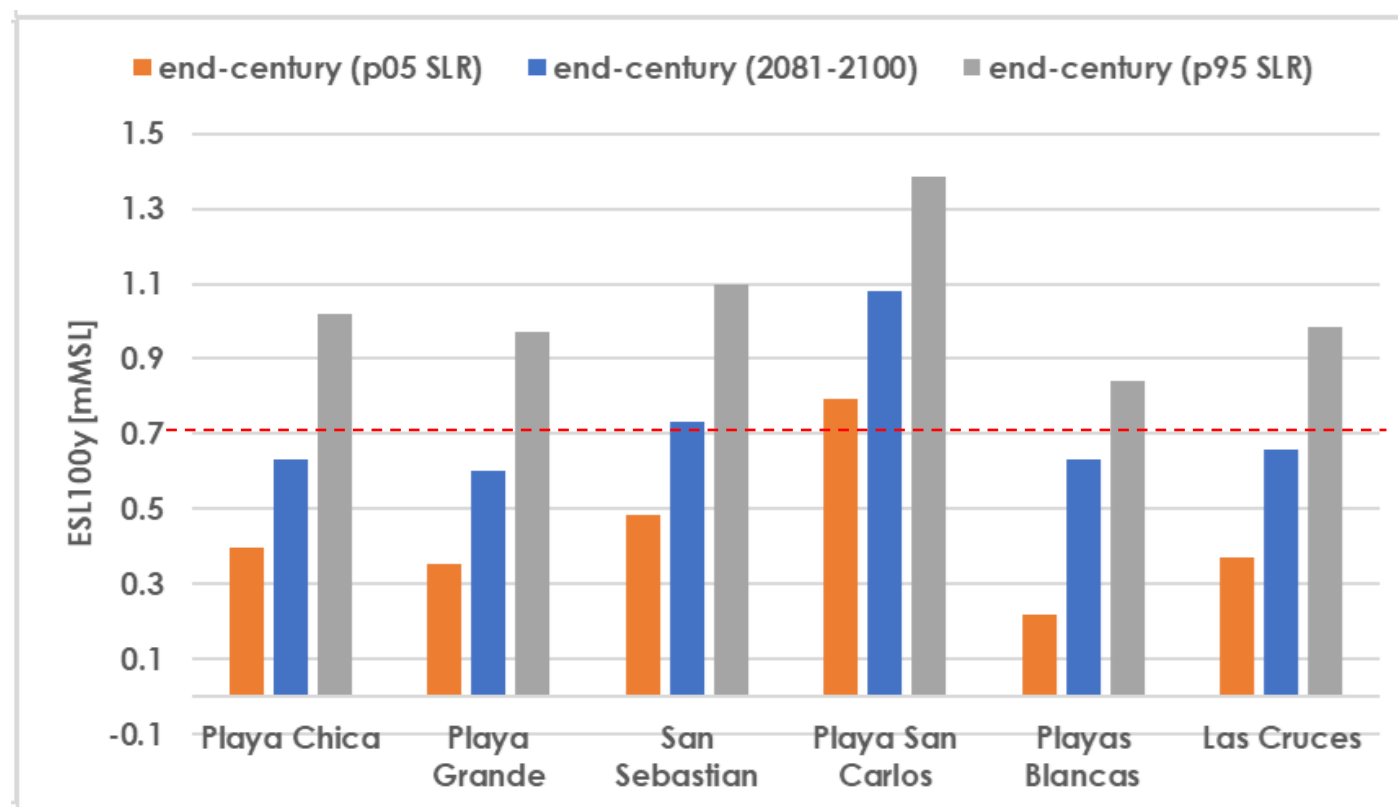
SEA LEVEL RISE

21 Climate Models (Hamburg University database)



Mean sea level rise projection with RCP 8.5						
Sea level rise process	2045			2100		
	5% (optimistic)	mean	95% (pessimistic)	5% (optimistic)	mean	95% (pessimistic)
Ocean Dynamics (Steric + dynamic + Inv. Barometer effect)	0.019	0.067	0.115	0.154	0.275	0.396
Glaciers	0.031	0.059	0.089	0.102	0.178	0.259
Greenland Dynamic	0.013	0.022	0.032	0.027	0.068	0.109
Antarctic Dynamic	0.002	0.018	0.033	-0.014	0.057	0.128
Greenland SMB	0.007	0.016	0.035	0.048	0.116	0.266
Groundwater depletion	0.003	0.016	0.028	-0.014	0.043	0.101
Antarctic SMB	-0.017	-0.008	-0.002	-0.086	-0.045	-0.016
GIA	-0.029	-0.012	0.005	-0.061	-0.025	0.011
Total SLR	0.100	0.177	0.261	0.399	0.667	0.998

100y RETURN LEVEL EVENT EXTREME SEA LEVELS AT CARTAGENA BAY

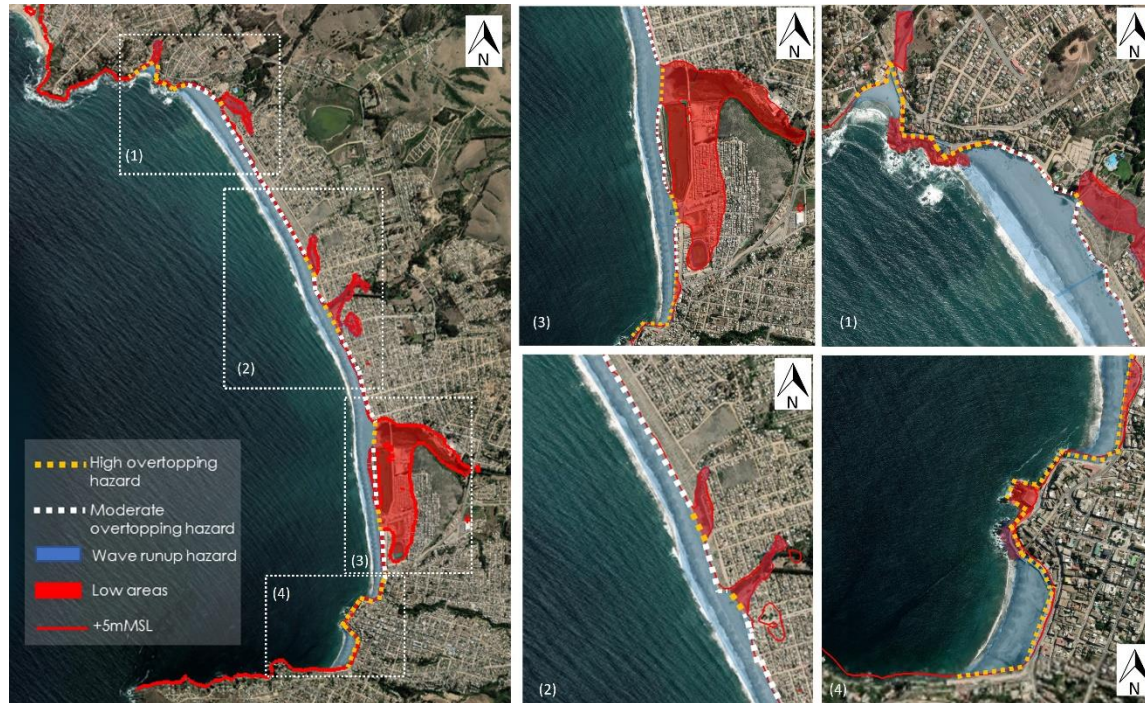


CONCLUSIONES

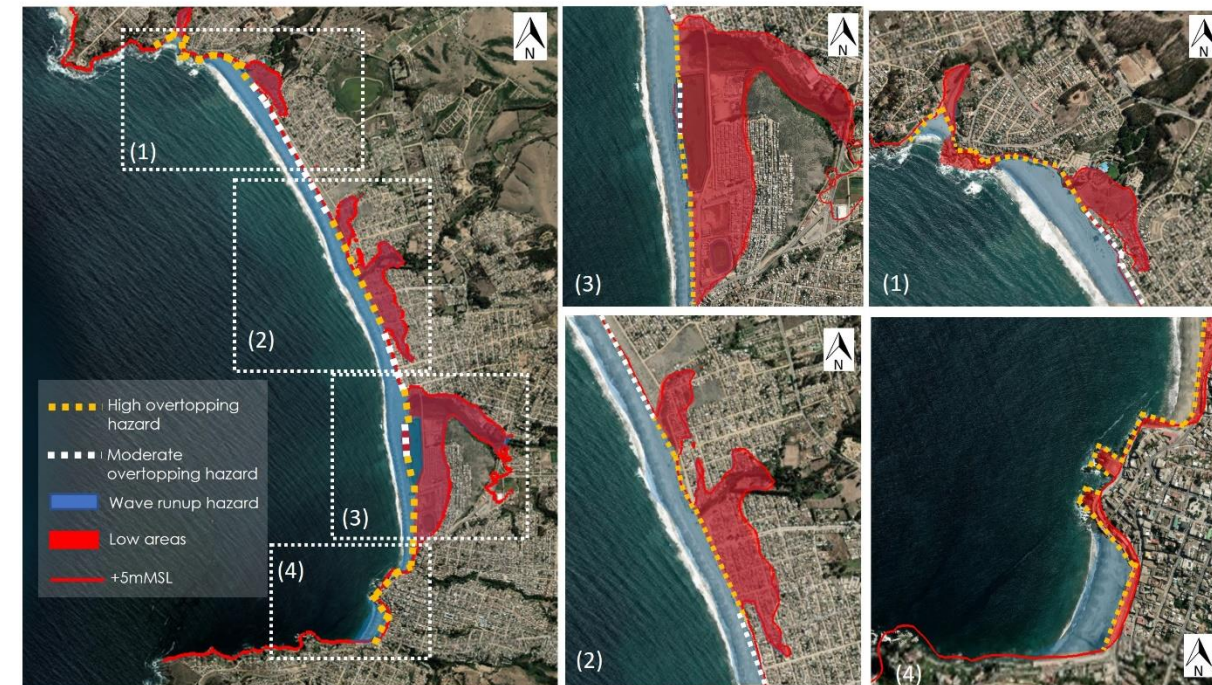
- Hacia finales de siglo, la zona de generación de olas de tormenta se desplazará gradualmente hacia latitudes más altas debido a un desplazamiento hacia el polo del cinturón de tormentas (de acuerdo con Hemer et al. 2013, Lobeto et al. 2021 y Morim et al. 2021). Esto inducirá olas de oleaje más largas de 0,5 a 1s.
- Además, la dirección de la onda media tenderá a desplazarse en sentido contrario a las agujas del reloj, con una rotación de unos 3°.
- Estos resultados sugieren la presencia de un "bloqueo" atmosférico a los sistemas de bajas presiones en latitudes más bajas, que puede explicarse por la migración hacia el sur del Anticiclón Subtropical del Pacífico (SPSA) (De acuerdo con Rykaczewski, 2015).
- En los sectores protegidos (parte norte y sur de la bahía de Cartagena). La amplificación prevista de la ESL será consecuencia del SLR, que podría superar los 60 cm a finales de siglo
- En las partes centrales de la bahía de Cartagena, el aumento previsto del ESL hacia finales de siglo estará por encima del SLR. Esto se debe principalmente al aumento del periodo de oleaje; sin embargo, esto debe confirmarse con modelos basados en procesos y modelación física
- Otro aspecto importante que no se ha considerado en este trabajo es que la zona costera de Chile es tectónicamente activa, por lo que la cota de inundación debe estimarse bajo una perspectiva de riesgos múltiples.

ESL vs COTA 10mMSL

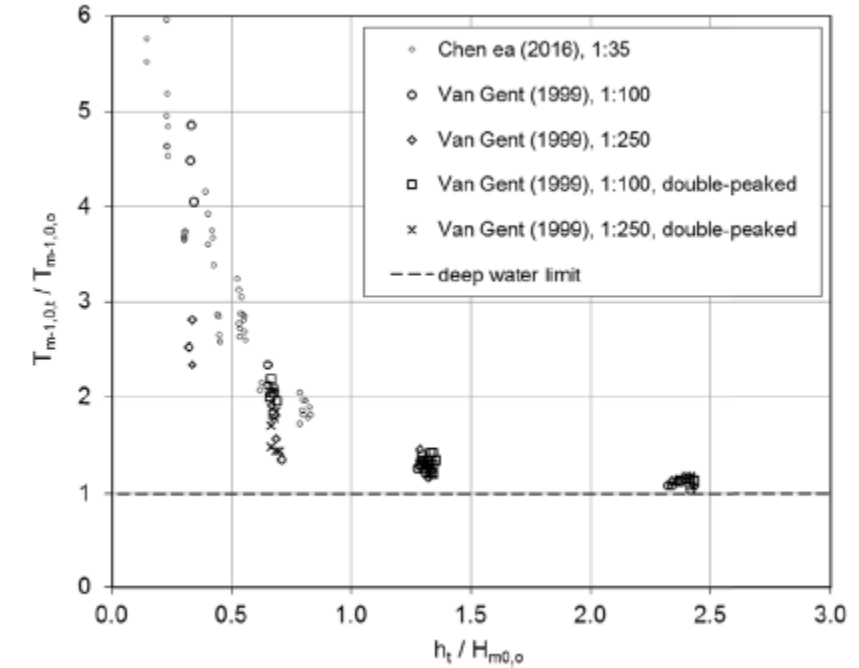
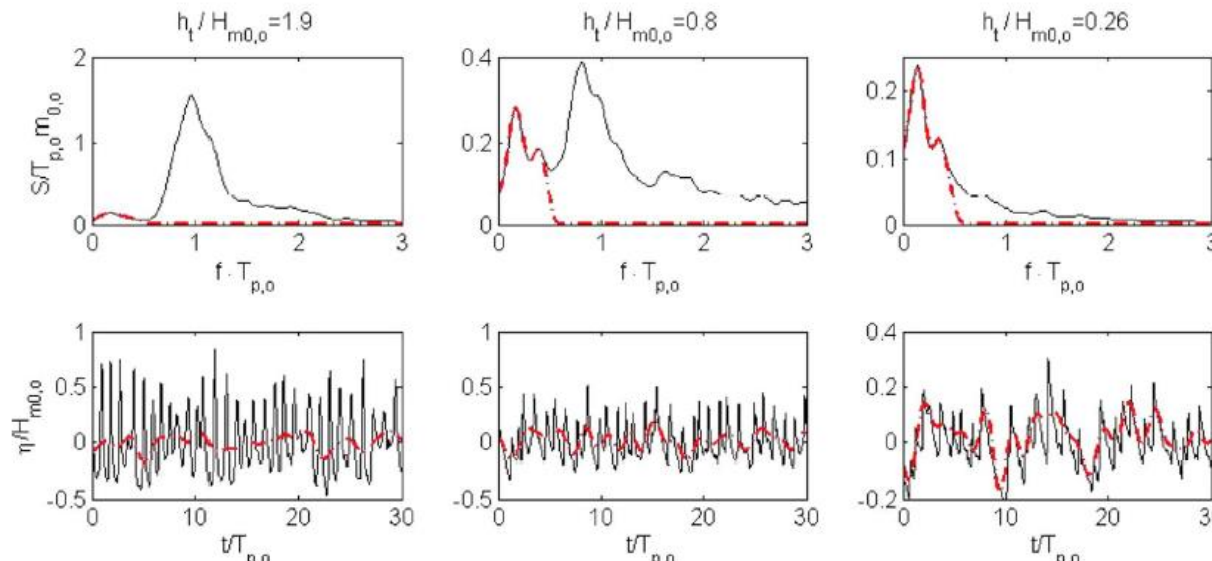
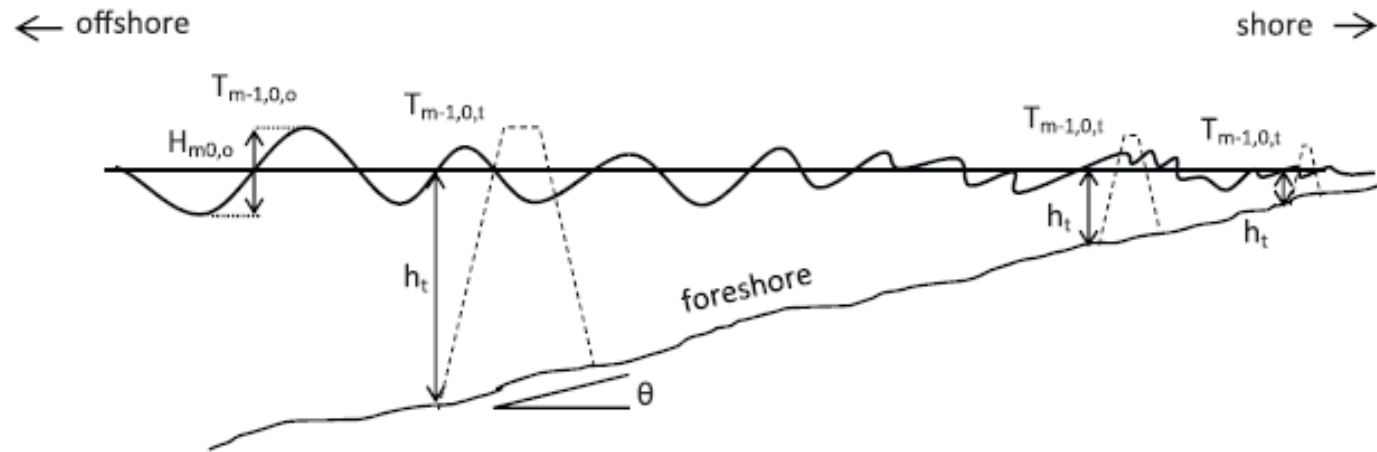
1985-2004



2081-2100



NEXT STEP: PHYSICAL MODELLING (MY PHD ☺)



Prof. Tom Bruce



Prof. Rodrigo Cienfuegos

¡Gracias!